# Effects of sky radiation on surface reflectance

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Surface and atmospheric radiative transfer communities seem to live on different "flat" planets.



The atmospheric community knows that there is something solid below the atmosphere which reflects sun light in a more complicated way than a simple Lambertian surface.

Lambertian "Equivalent" Reflectance (LER) is still widely used, however without a clear mathematical definition.

Assuming the surface BRF is known, how to define a LER?



### THE ROLE OF SKY RADIATIONS IN SHAPING SURFACE REFLECTANCE ANISOTROPY







### ATMOSPHERE-FREE BRF





### TOP-OF-ATMOSPHERE BRF

# BOTTOM-OF-ATMOSPHERE BRF

The **B**ottom-**O**f-**A**tmoisphere (BOA) BRF should not be confused with the atmospheric-free surface BRF.

Atmospheric correction methods usually do not make clear whether the BOA BRF or **A**tmospheric-**F**ree (AF) surface BRF is provided.

These differences can have a significant impact for lower atmospheric composition retrieval over land surfaces.



### SRF – ATM INTERACTIONS 1D VEGETATED SURFACE RTM

Simulation performed with the model of Gobron, et al. 1996. "A Semi-Discrete Model for the Scattering of Light by Vegetation." Journal of Geophysical Research 102: 9431–46.



















Aerosol optical thickness can change because:

- The particle concentration or extinction coefficient change in time
- Changes in the wavelength of (hyperspectral) observations









### SRF – ATM INTERACTIONS Constant instrinsic surface properties in the 320 – 750nm range





### SRF – ATM INTERACTIONS Let's assume constant intrinsic surface properties in the 320 – 750nm range





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BOA BRF IN THE PRINCIPAL PLANE IN THE 320 - 750nm SPECTRAL RANGE













### SRF – ATM INTERACTIONS Analysis of the AF – BOA BRF











WAVELENGTH

0.6

0.7

0.5

0.4

### NOW WITH CHANGING SURFACE CONDITIONS















## PRACTICAL SENTINEL-4 EXAMPLE



## PARIS EXAMPLE











### FROM AF BRF TO LER Method 1: LER = AF BRF

### $LER(t) = f(\Omega_s(t), \Omega_v; p_i)$



This approximation could be used in case of low optical thickness, *e.g.*, in the NIR spectral region. Provides the exact solution for the single scattering.



## FROM AF BRF TO LER Method 2: LER = AF DHR (black sky)

### $LER(t) = DHR(\Omega_s(t); p_i)$



This approximation is pretty meaningless as it does not respect the S4 illumination of viewing conditions.

## **FROM AF BRF TO LER** Method 2: LER - AF DHR





This approx nation is pretty meaningless as it does not resp. t the S4 illumination of viewing conditions.



## FROM AF BRF TO LER Method 3: LER = Diffuse BHR (white sky)

### LER $=BHR(p_i)$



This approximation could be used in case of high scattering optical thickness, e.g., in the UV region.



## FROM AF BRF TO LER Method 4: LER = BOA BRF LER(t, $\tau$ ) =g( $\Omega_s(t), \Omega_v; p_{\mu}\tau(t)$ )



This approximation is the best choice but requires the knowledge of  $\tau(t)$  and the irradiance field at the surface.





## CONCLUSIONS (1)

Sky radiation plays an important role in shaping surface BRF;

- BOA BRF and atmospheric free BRF are often mixed though they might exhibit important differences
  - In the blue-NIR spectral region;
  - When the scattering optical thickness is important;
  - In the hot spot conditions;

These differences might have an important impact for lower atmosphere trace gas retrieval over land surfaces.

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